

# 1 Derive continuity equation for regolith thickness

$$H(x, y, t) = H_0(x, y) + U(x, y, t) + V(x, y, t, H) \quad (1)$$

Where  $H_0$  is the initial topography,

$U$  is the vertical component of tectonic displacement,

and  $V$  is the geomorphic displacement.

Because  $U$  is the same for  $H$  and  $B$  at a given point  $(x, y)$ , it does not change the regolith thickness:

$$h(x, y) = (H(x, y) - U(x, y)) - (B(x, y) - U(x, y))$$

or

$$h(x, y) = H(x, y) - B(x, y)$$

Therefore, it is convenient to consider soil continuity in terms of soil thickness.

We use the derivation of conservation of transportable (regolith) for our control area (note that is a 2 dimensional problem):

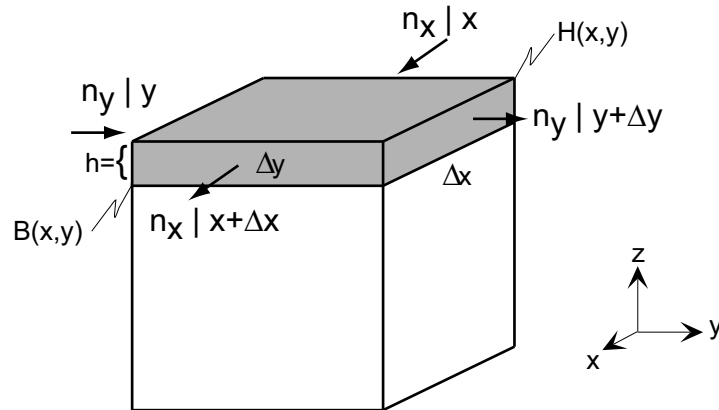


Figure 1: regolith continuity for a hillslope element

$$\begin{aligned}
& \left\{ \text{net rate of mass efflux from control area} \right\} + \\
& \left\{ \text{net rate of mass accumulation in control area} \right\} - \\
& \left\{ \text{rate of production of mass within control area} \right\} \\
& = 0
\end{aligned}$$

So let's take each term and evaluate it:

1. Net rate of mass efflux from control area. Consider the mass flux vector *per unit width*

$$\vec{n} = \rho \vec{v} h = \rho_s \vec{q}_s \quad [ML^{-3} \cdot L^1 T^{-1} \cdot L^1] = [ML^{-3} \cdot L^2 T^{-1}] = ML^{-1} T^{-1} \quad (2)$$

where  $h$  is the thickness of the regolith or mobile layer (and through which the flux moves)  $[L]$

$\vec{v}$  is the sediment velocity  $[L]$

$\rho_s$  is sediment density  $[ML^{-3}]$

$\vec{q}_s$  is sediment volume flux *per unit width*  $[L^2 T^{-1}]$

If we resolve the mass flux vector into its cartesian components, we have:

in the  $x$  direction:  $n_x \Delta y|_{x+\Delta x} - n_x \Delta y|_x$

in the  $y$  direction:  $n_y \Delta x|_{y+\Delta y} - n_y \Delta x|_y$

in the  $z$  direction:  $n_z = 0$

2. Net rate of accumulation of mass in the control area:

$$\rho_s \frac{\partial h}{\partial t} \Delta x \Delta y \quad (3)$$

3. Net rate of production of mass in the control area:

$$-\rho_r \frac{\partial B}{\partial t} \Delta x \Delta y \quad \rho_r \text{ is the rock density} \quad (4)$$

The negative sign ensures that a negative change in  $B$  causes positive mass change.

Put it all together:

$$\begin{aligned}
& \left\{ (n_x \Delta y|_{x+\Delta x} - n_x \Delta y|_x) + (n_y \Delta x|_{y+\Delta y} - n_y \Delta x|_y) \right\} + \\
& \rho_s \frac{\partial h}{\partial t} \Delta x \Delta y - \\
& -\rho_r \frac{\partial B}{\partial t} \Delta x \Delta y = 0
\end{aligned}$$

Divide through by  $\Delta x \Delta y$  and simplify

$$\begin{aligned}
\left(\frac{\Delta n_x}{\Delta x} + \frac{\Delta n_y}{\Delta y}\right) + \rho_s \frac{\partial h}{\partial t} + \rho_r \frac{\partial B}{\partial t} &= 0 \\
\left(\frac{\partial}{\partial x} n_x + \frac{\partial}{\partial y} n_y\right) + \rho_s \frac{\partial h}{\partial t} + \rho_r \frac{\partial B}{\partial t} &= 0 \\
\nabla \cdot \vec{n} + \rho_s \frac{\partial h}{\partial t} + \rho_r \frac{\partial B}{\partial t} &= 0 \\
\rho_s \nabla \cdot \vec{q}_s + \rho_s \frac{\partial h}{\partial t} + \rho_r \frac{\partial B}{\partial t} &= 0
\end{aligned}$$

$$\rho_s \frac{\partial h}{\partial t} = -\rho_s \nabla \cdot \vec{q}_s - \rho_r \frac{\partial B}{\partial t} \quad (5)$$

Thus (5) is the continuity equation for regolith thickness.

If we want to go for transport limited conditions,  $\rho_r \frac{\partial B}{\partial t} = 0$ , and you can divide through by  $\rho_s$ :

$$\frac{\partial h}{\partial t} = \frac{\partial H}{\partial t} = -\nabla \cdot \vec{q}_s$$

or in 1 dimension:

$$\frac{\Delta H}{\Delta t} = -\frac{\Delta \vec{q}_s}{\Delta x}$$

## 2 Evaluation of dimensions

If we look again at the first portion of (5),

$$\begin{aligned} \left(\frac{\Delta n_x}{\Delta x} + \frac{\Delta n_y}{\Delta y}\right) + \rho_s \frac{\partial h}{\partial t} + \rho_r \frac{\partial B}{\partial t} &= 0 \\ \rho_s \left(\frac{\Delta q_{sx}}{\Delta x} + \frac{\Delta q_{sy}}{\Delta y}\right) + \rho_s \frac{\partial h}{\partial t} + \rho_r \frac{\partial B}{\partial t} &= 0 \end{aligned}$$

Focusing on the first term:

$$\rho_s \left(\frac{\Delta q_{sx}}{\Delta x} + \frac{\Delta q_{sy}}{\Delta y}\right) = \rho_s \nabla \cdot \vec{q}_s$$

Dimensionally:

$$[ML^{-3}] \left( \frac{[L^2 T^{-1}]}{[L]} + \frac{[L^2 T^{-1}]}{[L]} \right)$$

$$[ML^{-3}] ([L^1 T^{-1}] + [L^1 T^{-1}])$$

$$[ML^{-3}] [L^1 T^{-1}] = \rho_s \nabla \cdot \vec{q}_s$$

So, indeed

$$[L^1 T^{-1}] = \nabla \cdot \vec{q}_s$$